

An overview of the state-of-the-art in meteotsunami research and potential operational systems

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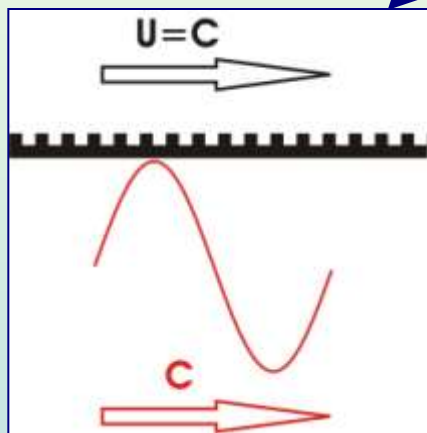
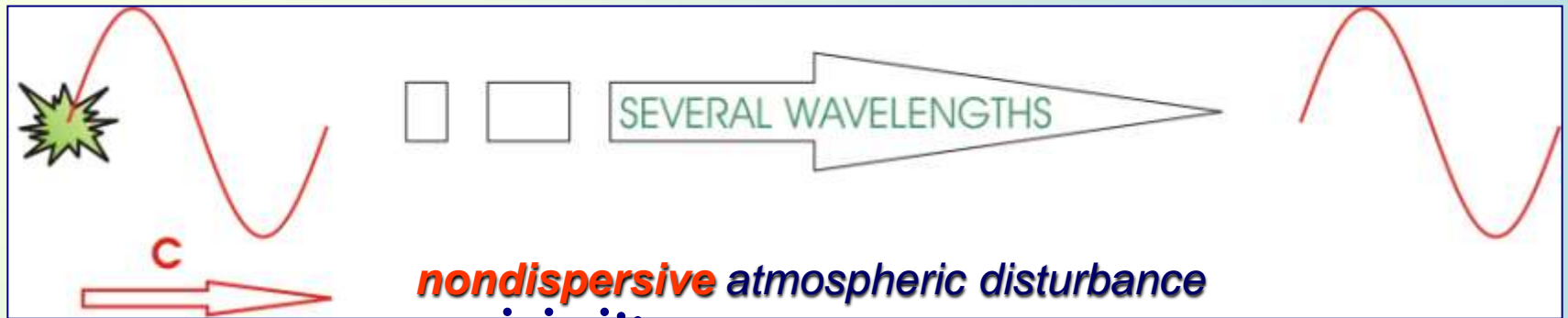
- Present knowledge on the phenomenon
 - The source
 - Air-to-ocean energy transfer (external resonance)
 - Coastal amplification (internal resonance)
- Common synoptic characteristics
- Operational possibilities



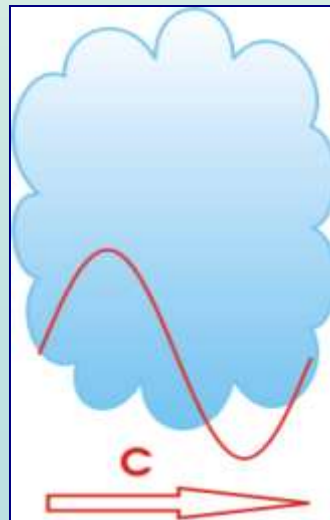
Meteotsunami, atmospherically induced destructive ocean waves in the tsunami frequency band, is multiresonant phenomenon which occur when having (Monserat *et al.*, NHESS, 2006):

- A **harbour** (bay, inlet or gulf) with definite **resonant** properties and high Q-factor.
- **Strong small-scale atmospheric disturbance** (a pressure jump or a train of internal atmospheric waves).
- **Propagation of the atmospheric disturbance toward the entrance** to the harbour.
- **External resonance** (Proudman, Greenspan or shelf resonance) between the atmospheric disturbance and ocean waves.
- **Internal resonance** between the dominant frequency of the arriving open-ocean waves and the fundamental harbour mode frequency.



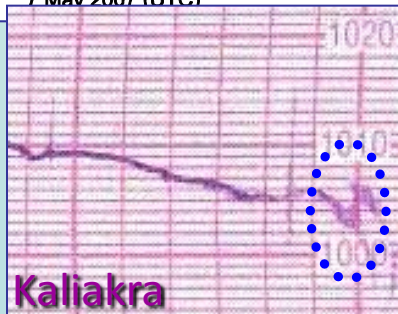
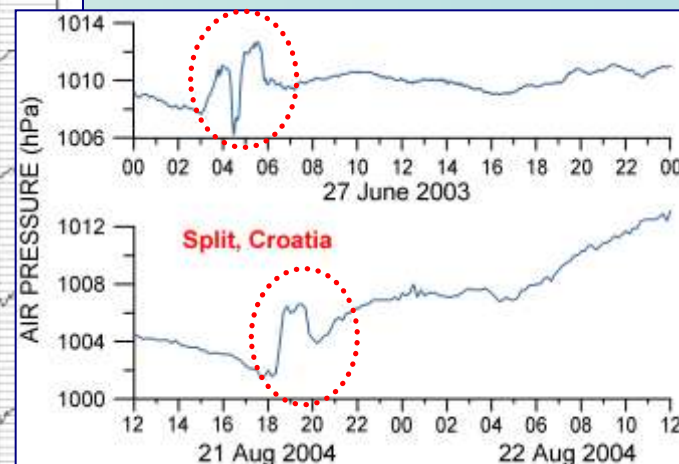
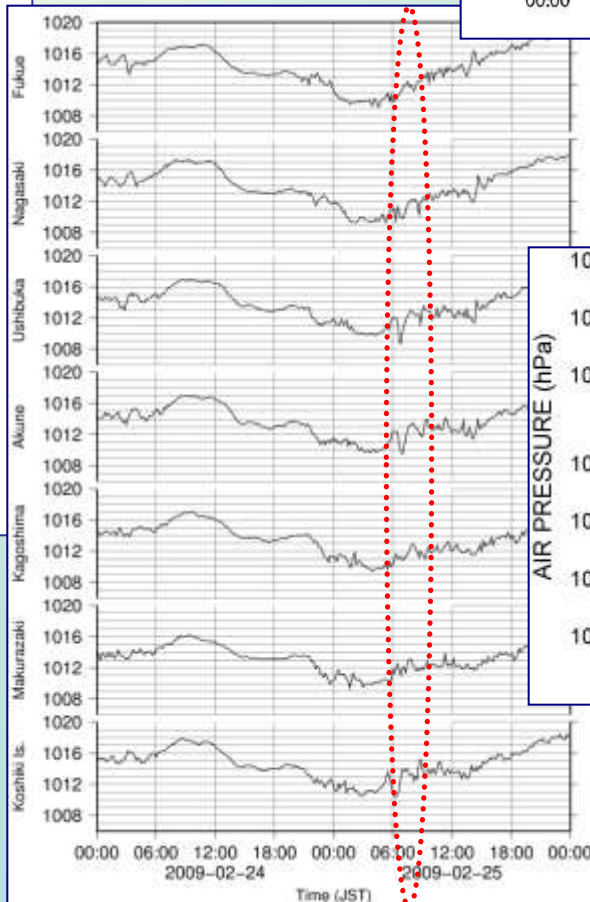
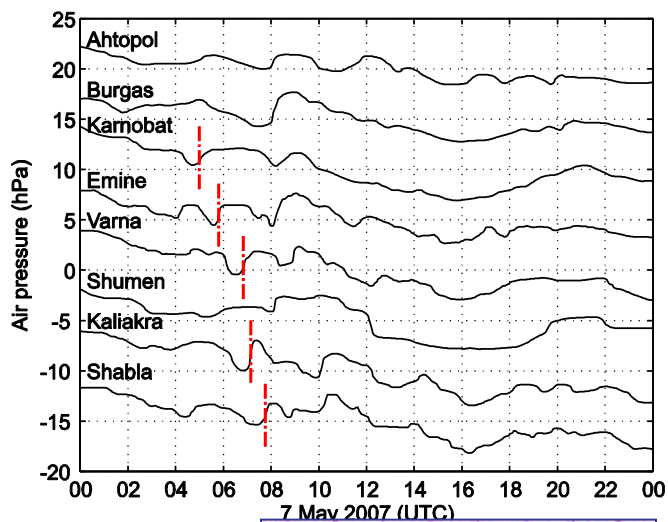
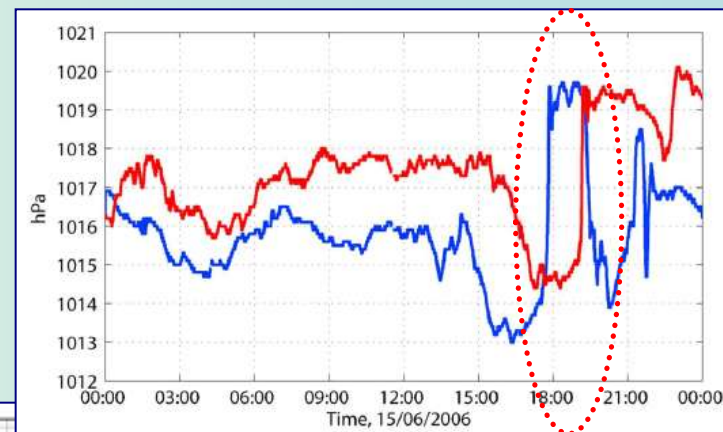


wave-CISK



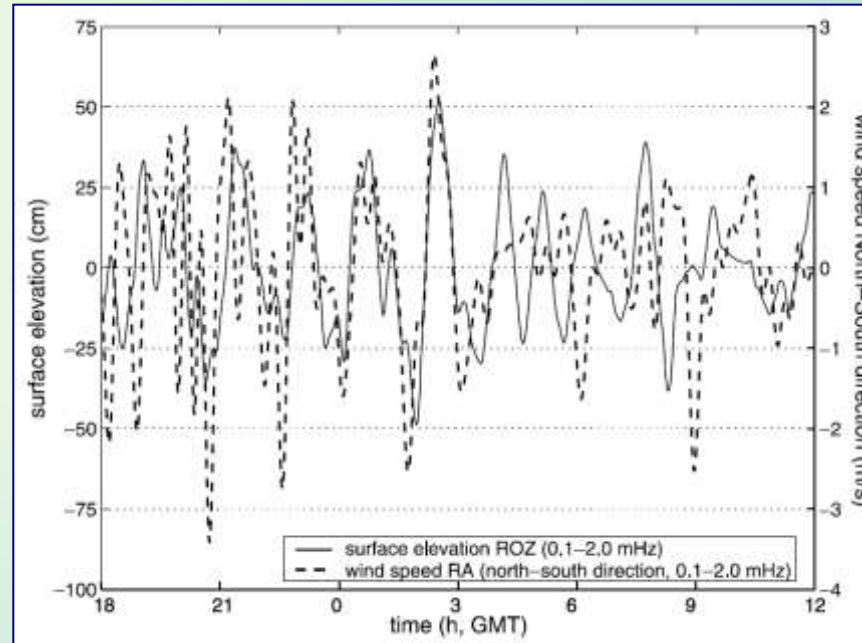
air pressure disturbance - examples

The source



Battjes and de Long, JGR-O, 2004:

- port of Rotterdam, North Sea
- cross-shore travelling wind disturbances

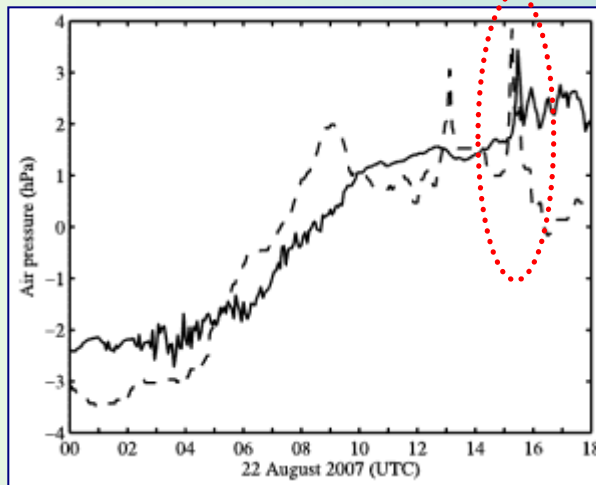
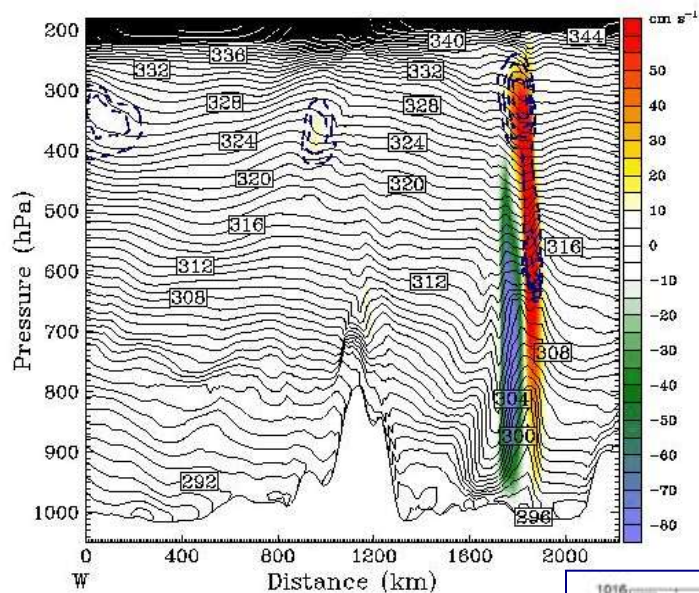


Wind gusts play a secondary role to explain the meteotsunami formation

- Vilibić et al., JGR-O, 2004
- Vilibić et al., JMS, 2005
- Renault et al., GRL, 2011

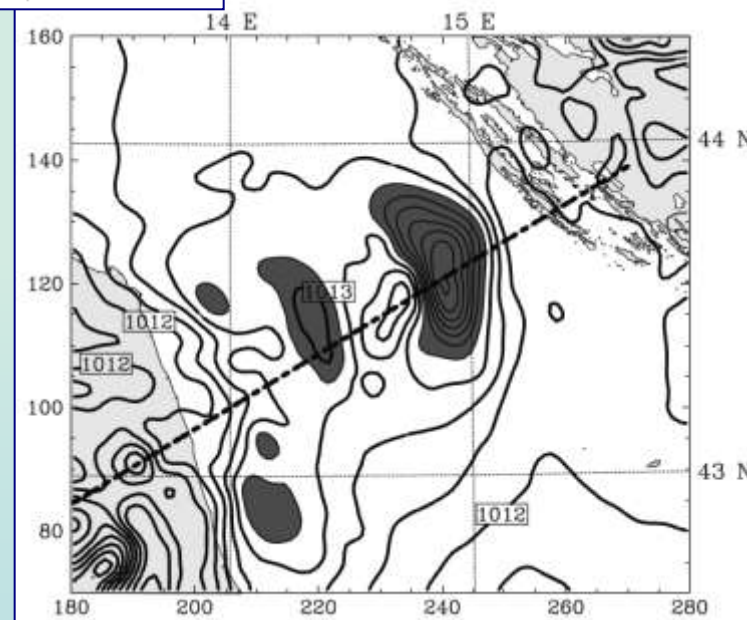
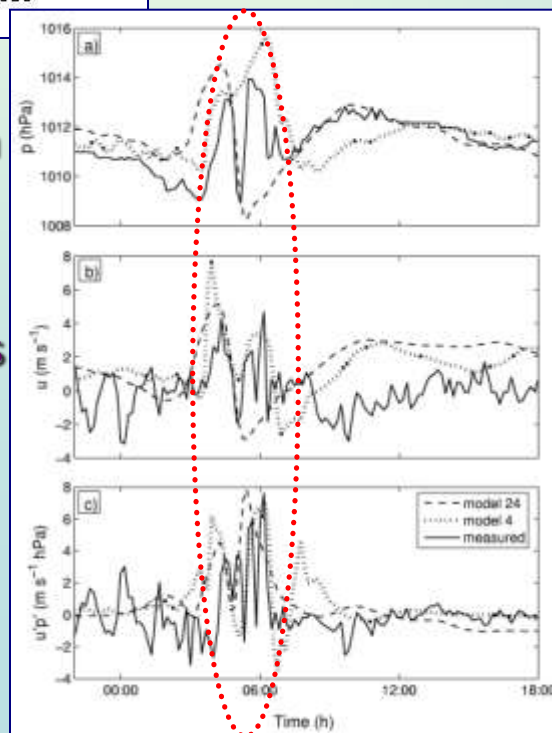


The source



- Šepić *et al.*, JGR-O, 2009:
- 22 August 2008, Adriatic
 - WRF mesoscale model
 - wave-duct, capped by convective instability
 - generated over mountains
 - width of a few tens of km
 - not so good verification

- Belušić *et al.*, JGR-A, 2007:
- 27 June 2007, Adriatic Sea
 - MM5 mesoscale model
 - wave-CISK within MCS
 - no wave-duct present
 - generated over mountains
 - not so good verification



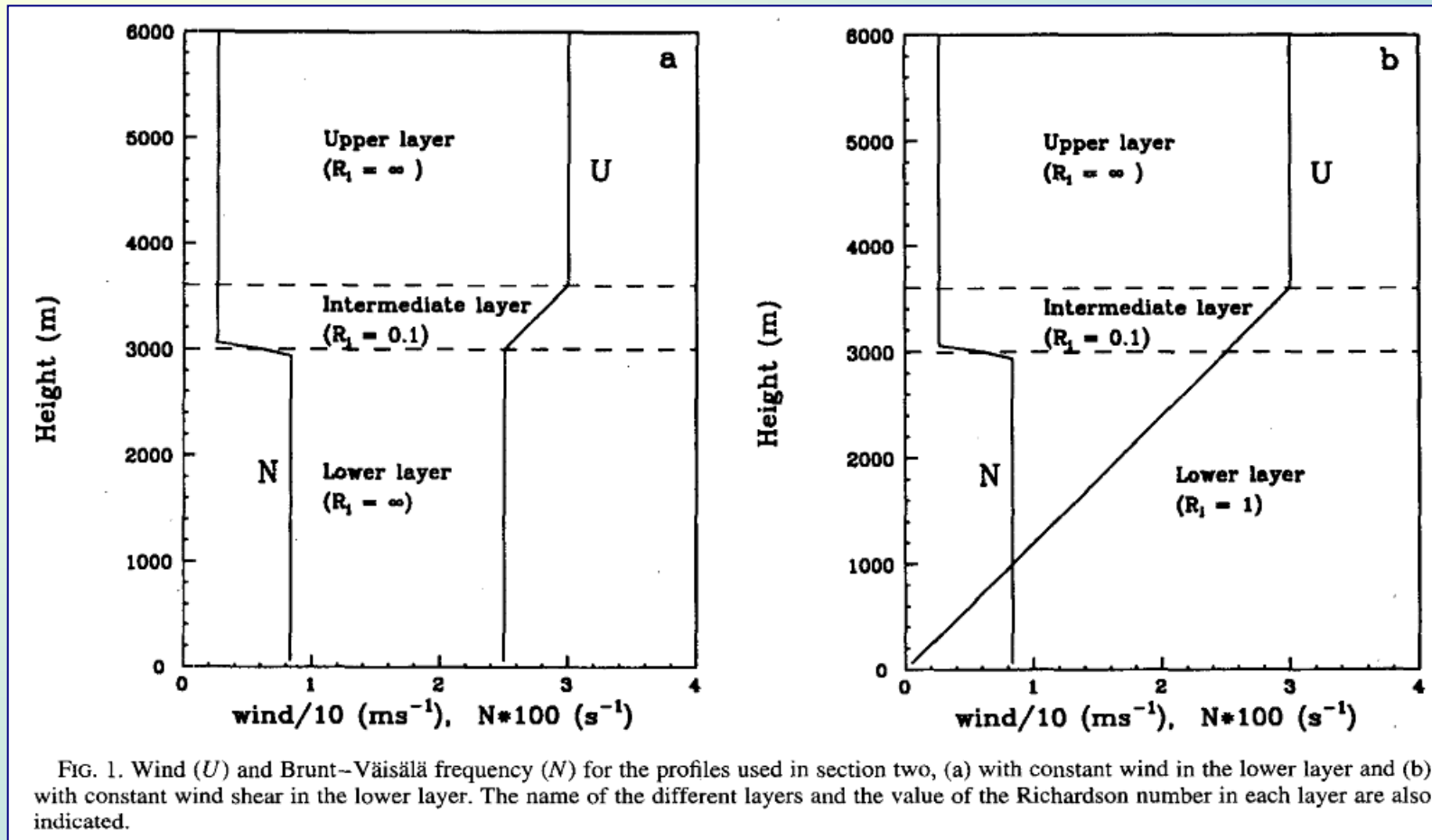


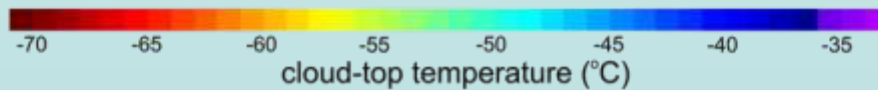
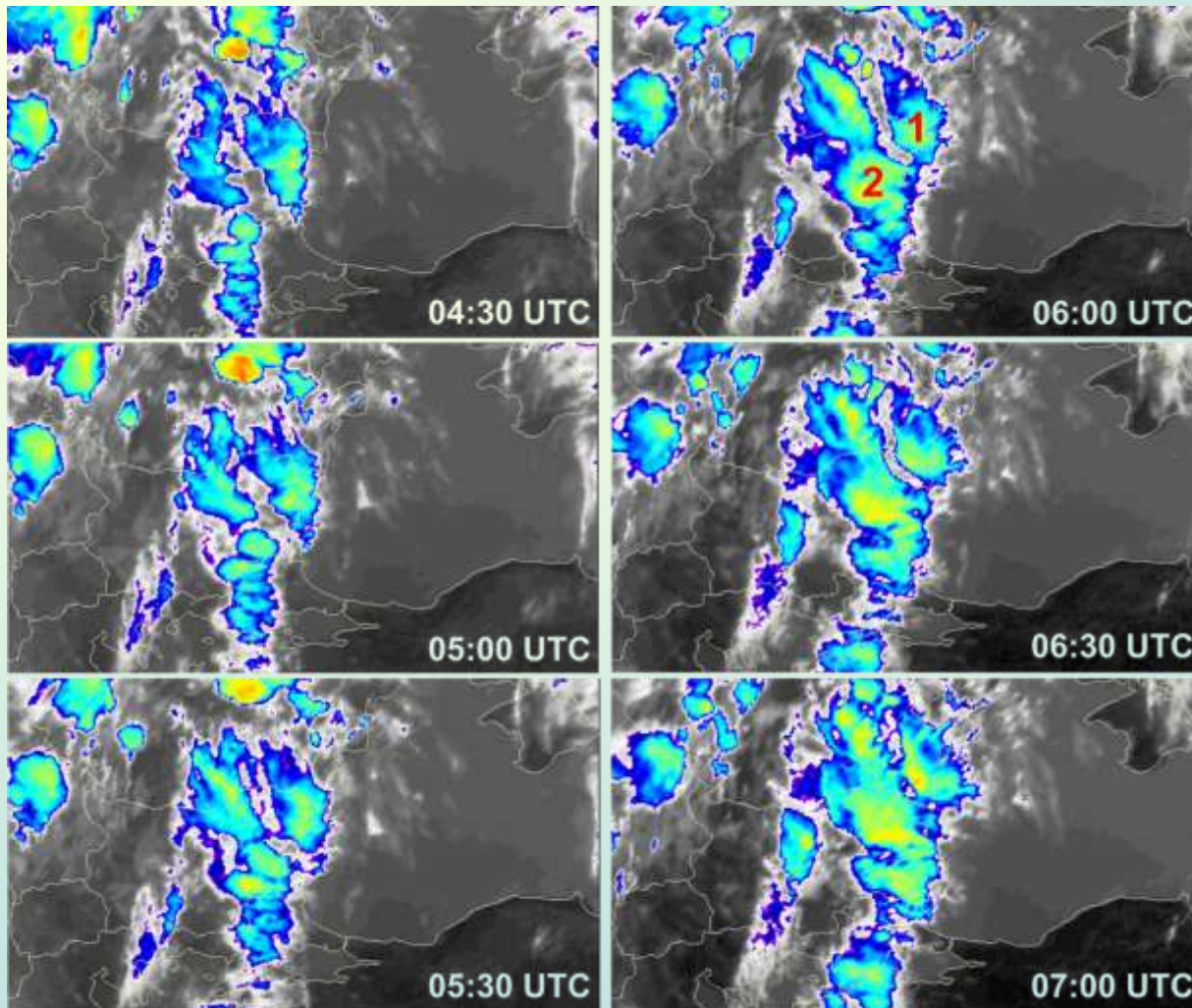
FIG. 1. Wind (U) and Brunt-Väisälä frequency (N) for the profiles used in section two, (a) with constant wind in the lower layer and (b) with constant wind shear in the lower layer. The name of the different layers and the value of the Richardson number in each layer are also indicated.

Monserrat and Thorpe, JAS, 1996:

- wave duct theory
- reflectance layer above stable layer
- wave speed equal to wind speed at reflectance layer
- wave speed dependable on the wind shear
- applied on the Balearic rissagas



The source



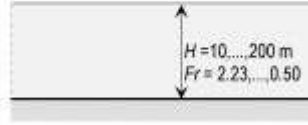
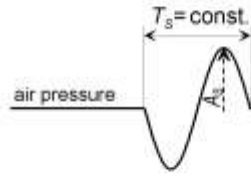
- Vilibić *et al.*, JGR-O, 2010:
- 7 May 2007, Black Sea
 - wave-duct, capped by a convective system
 - width of a few tens of km



External resonance

Proudman resonance

disturbance direction
 $U = \text{const.} = 22.15 \text{ m/s}$



$$Fr = \frac{U}{\sqrt{gH}}$$

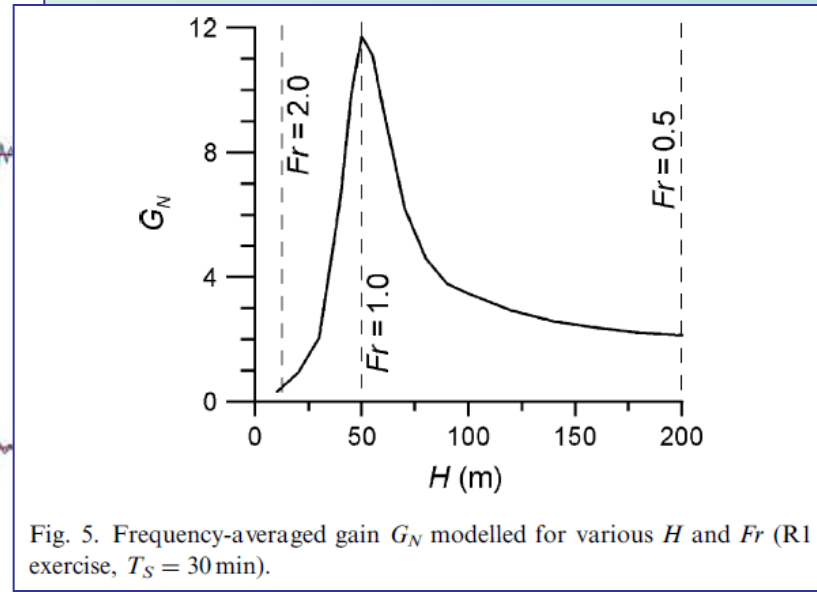
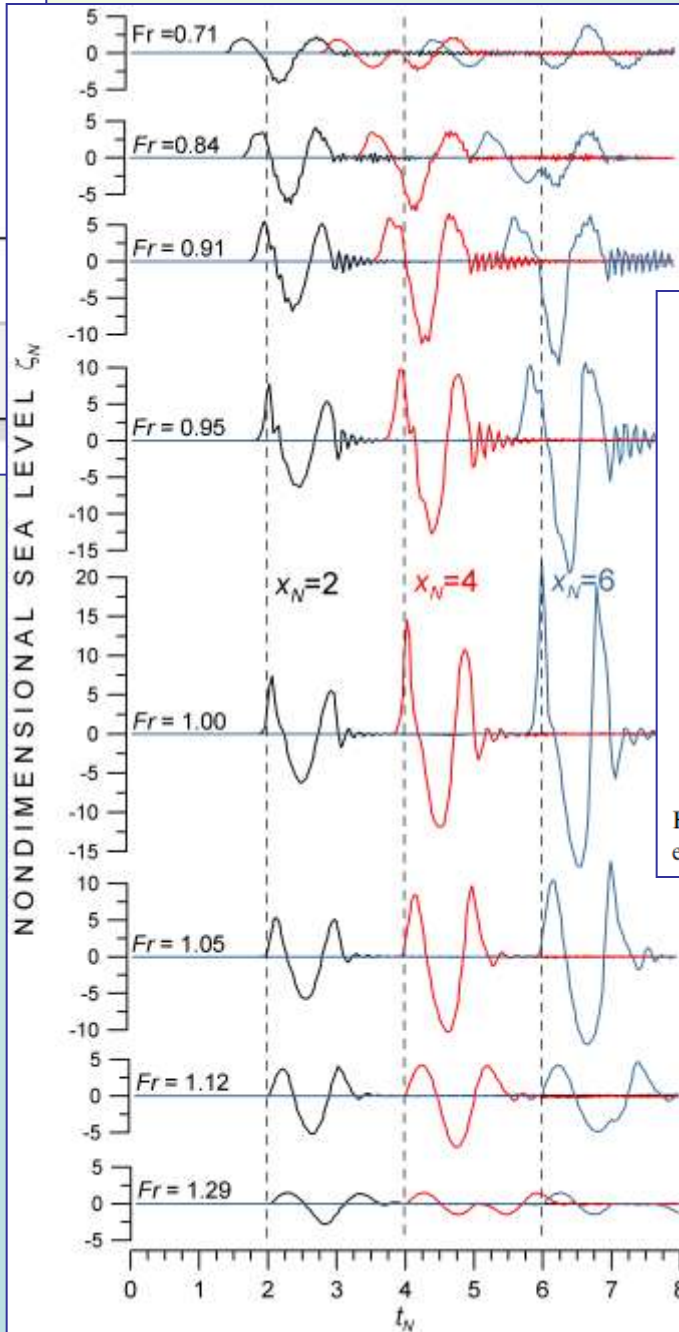


Fig. 5. Frequency-averaged gain G_N modelled for various H and Fr (R1 exercise, $T_S = 30 \text{ min}$).

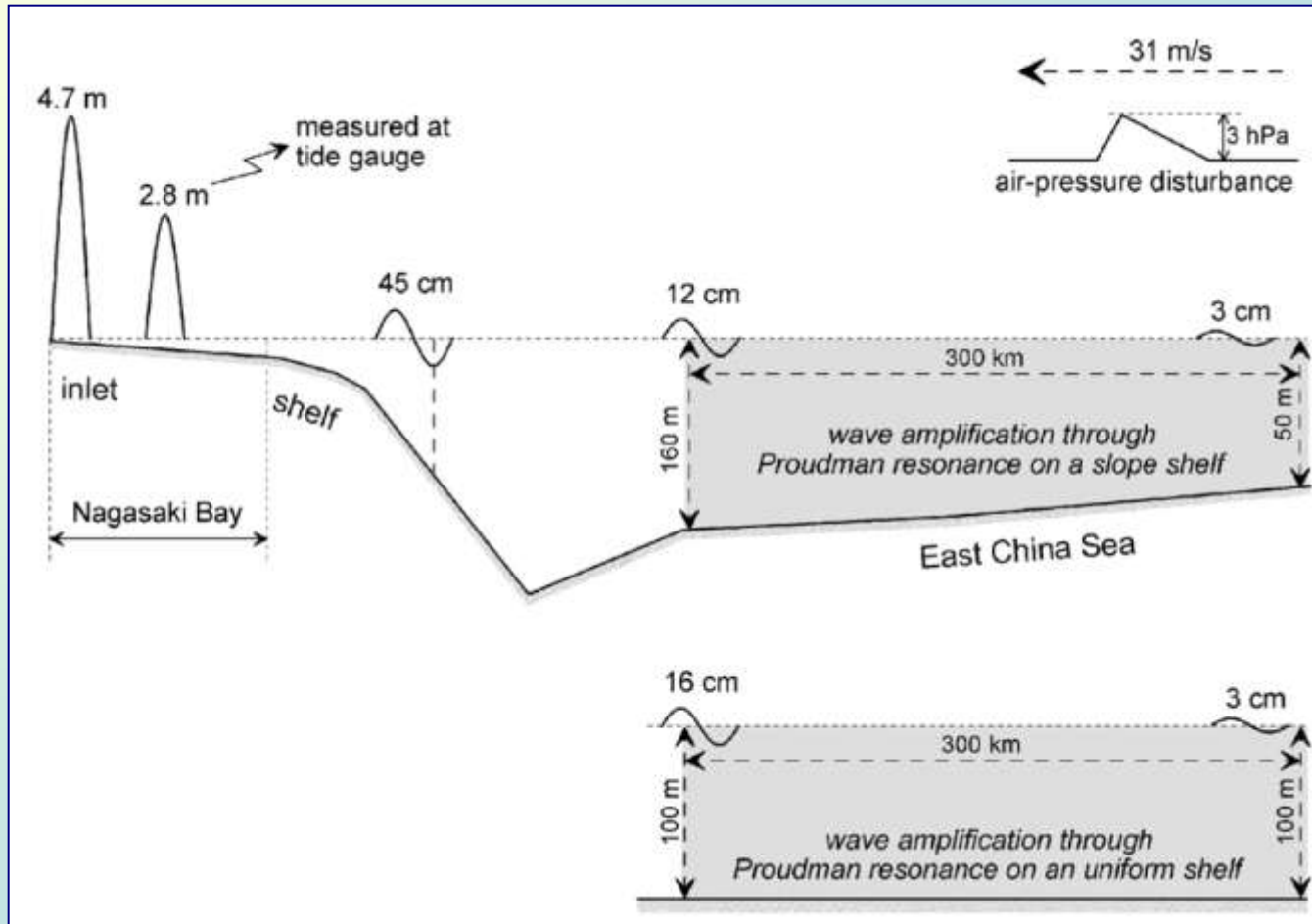
Vilibić, CSR, 2008:

- open ocean resonance quite sensitive to depth changes



External resonance

Proudman resonance



Vilibić, CSR, 2008:

- the presence of a slope is reducing the resonance efficiency



External resonance

Proudman resonance

Vilibić *et al.*, PAGEOPH, 2008

Orlić *et al.*, JGR-O, 2010:

- external resonance at a particular area is favoured by a range of atmospheric disturbance speeds and directions

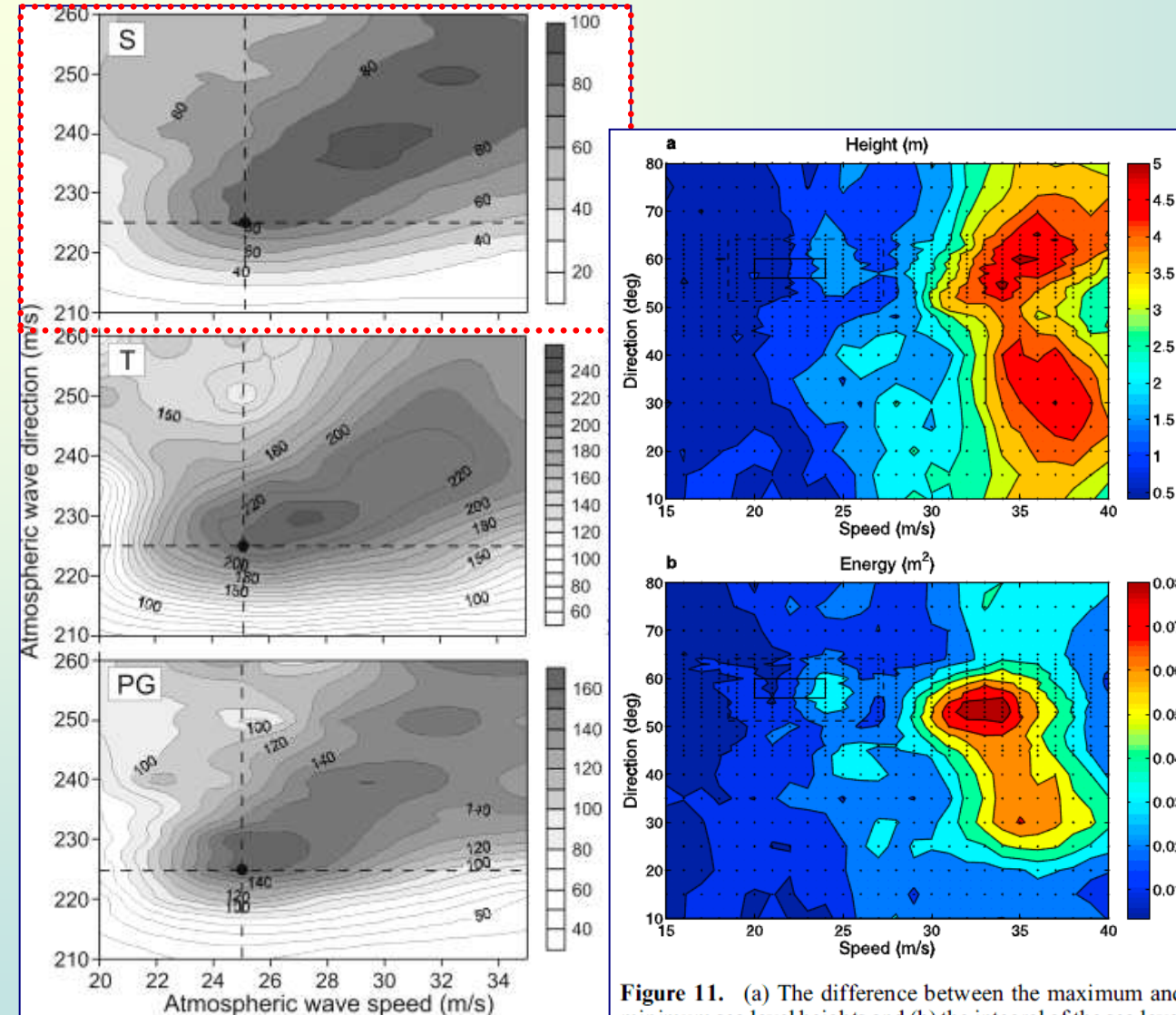
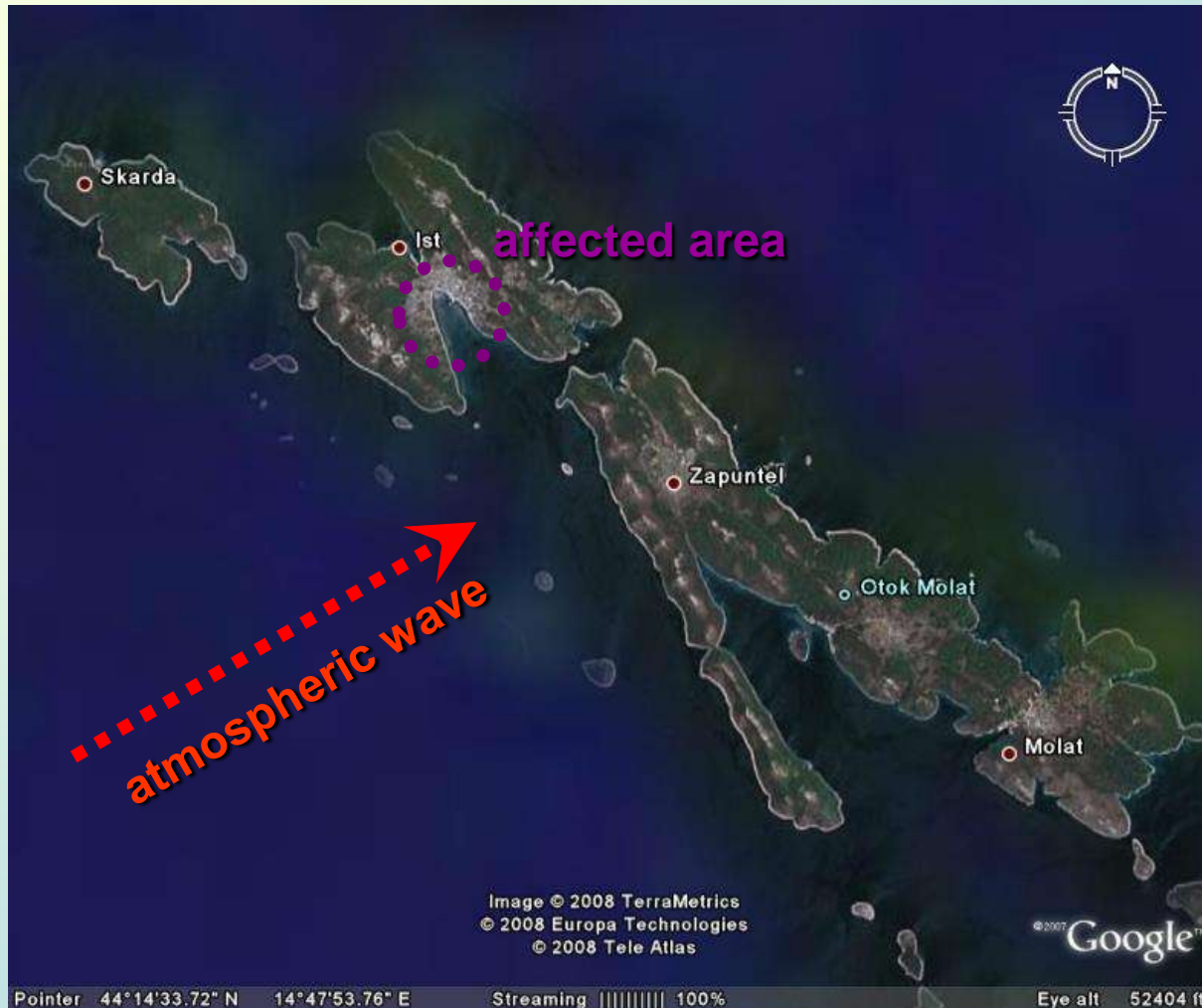


Figure 11. (a) The difference between the maximum and minimum sea level heights and (b) the integral of the sea level energy density spectrum obtained for the head of Vela Luka Bay by forcing the Adriatic model with the boxcar air pressure perturbation traveling at various speeds (c) and directions (γ). Each dot represents a run of the Adriatic model; the total number of runs is 754. The solid rectangle indicates the ranges of speeds and directions determined by Orlić [1980], and the dashed rectangle marks the ranges obtained in this paper by assuming that the timing error of barographs used in the Adriatic in 1978 was ± 10 min.



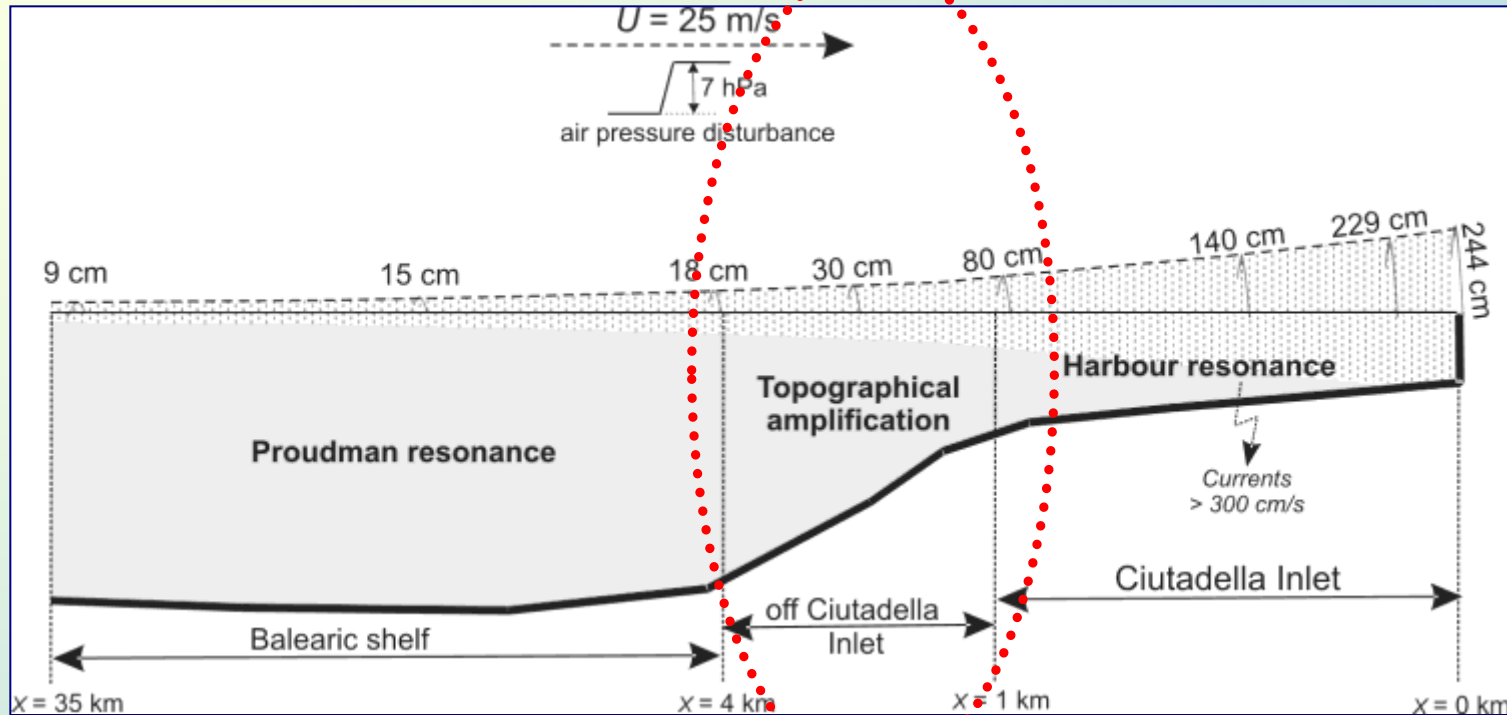


Propagation of the atmospheric disturbance toward the entrance to the harbour. → not always true

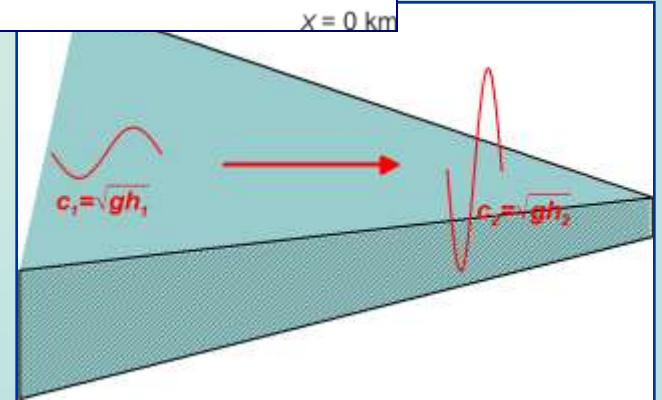
→ reflection?



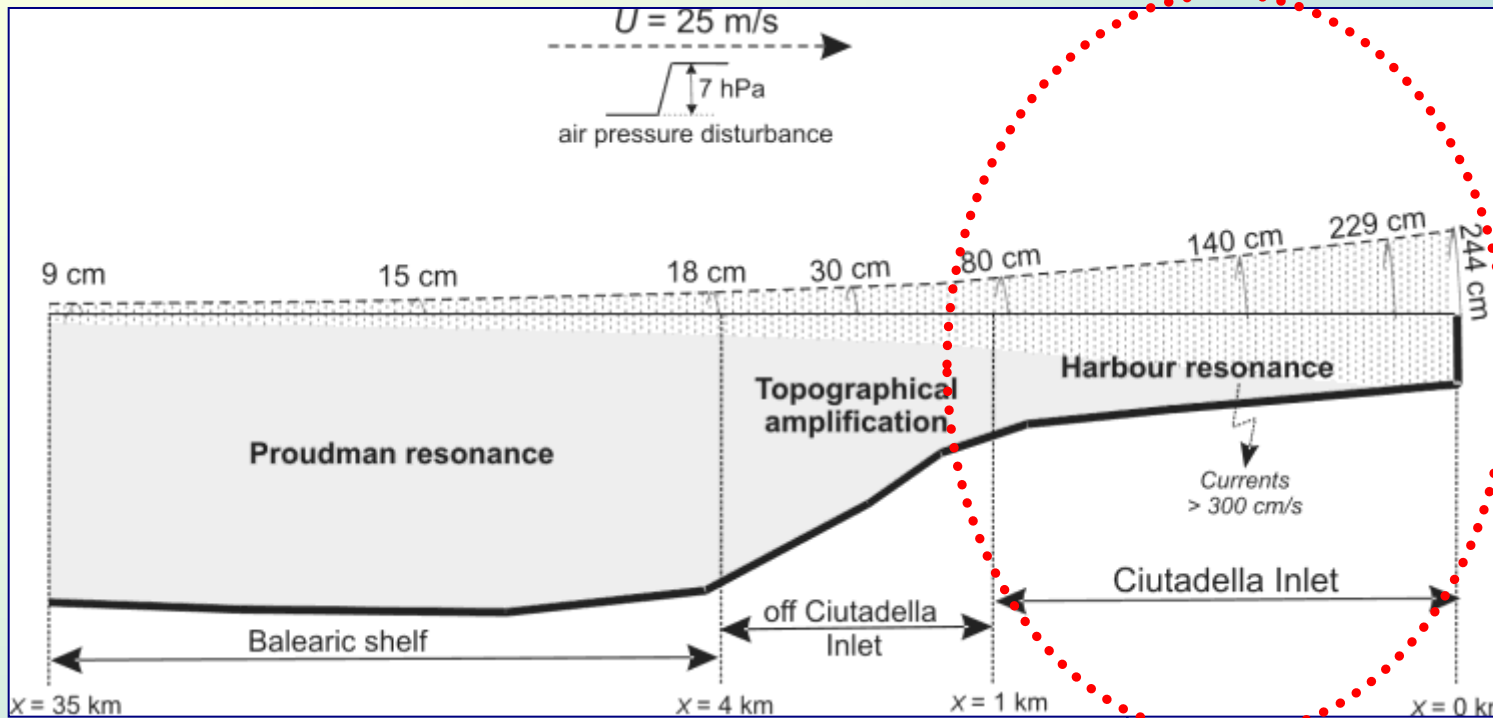
Topographical amplification



Vilibić *et al.*, PAGEOPH, 2008:
 - topographical amplification is
 as for any tsunami waves



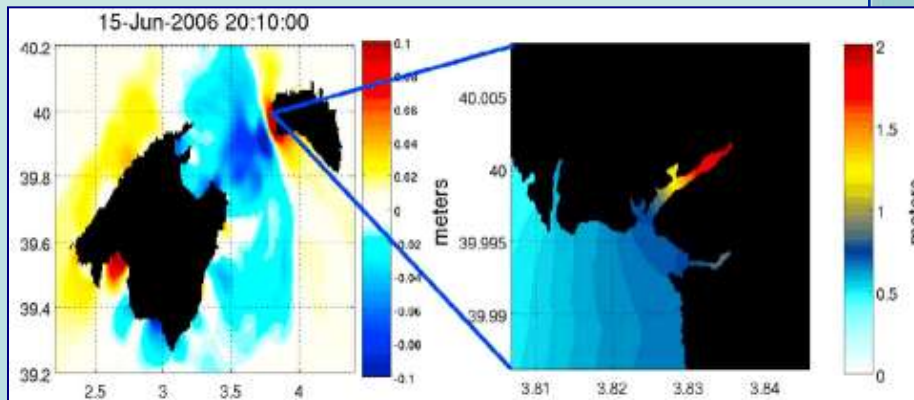
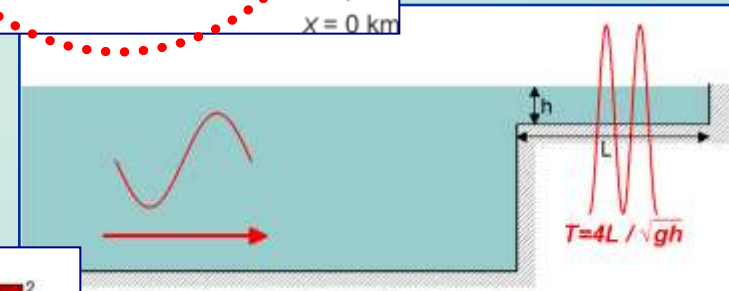
Harbour resonance



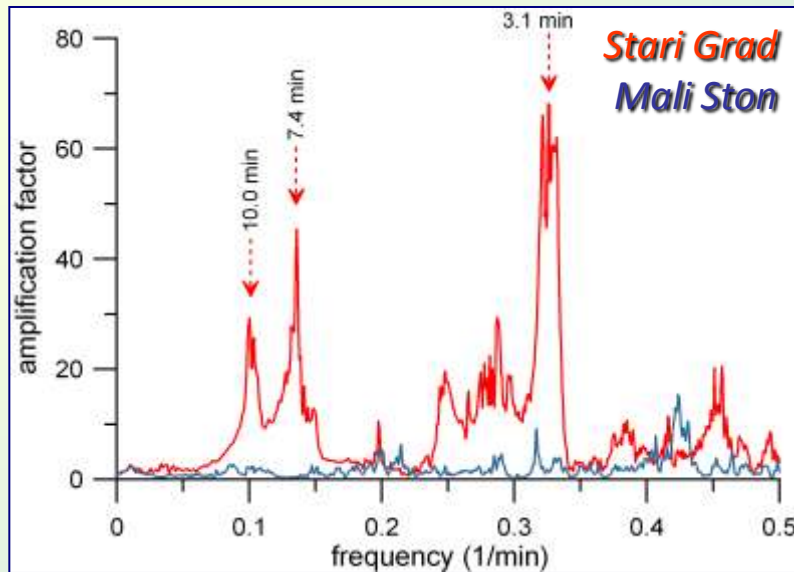
Vilibić *et al.*, PAGEOPH, 2008

Renault *et al.*, GRL, 2011

- harbour resonance is as for any tsunami waves



Harbour resonance



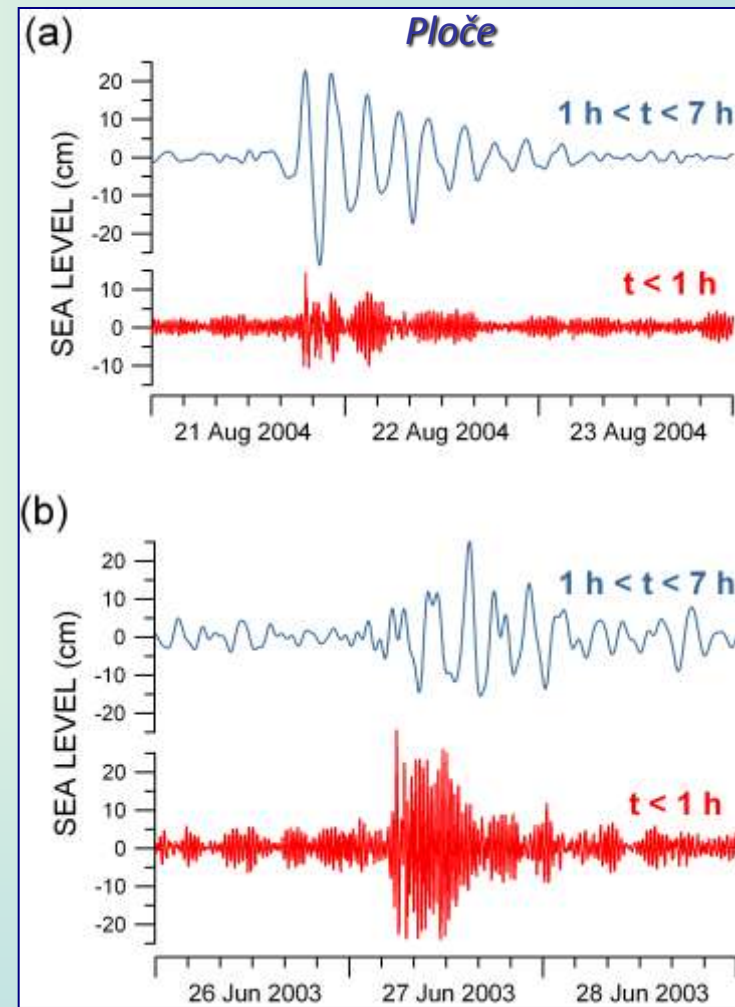
Vilibić and Šepić, PCE, 2009:
 - amplification factor vary with frequency

$$H^2(f) = \frac{1}{(1 - f/f_0)^2 + Q^{-2}(f/f_0)^2}$$

$$Q^{-1} = \frac{dE/dt}{\omega E} = 2\beta$$

$$Q = \frac{L}{l}, \text{ rectangular basin}$$

Rabinovich, 2009

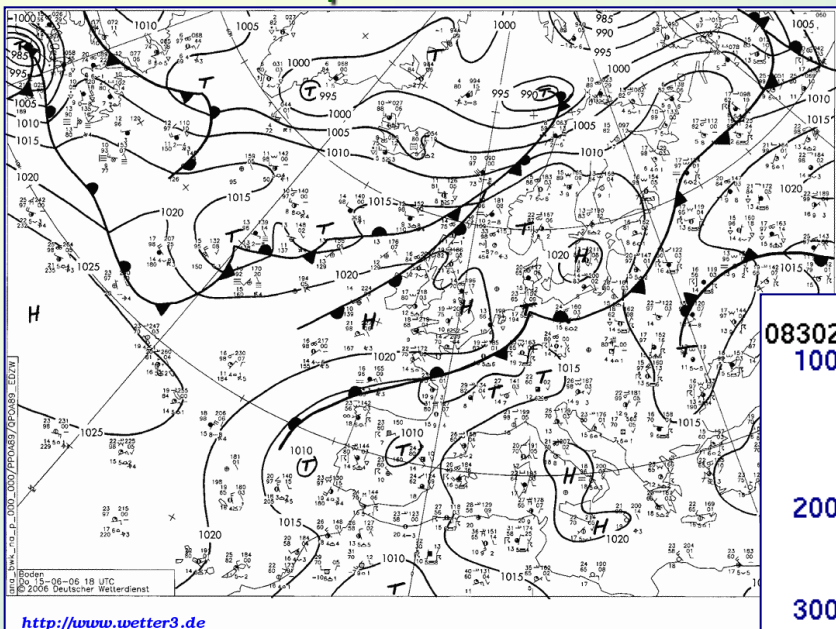


Vilibić, JMS, 2005:
 - energy of incoming waves do matter

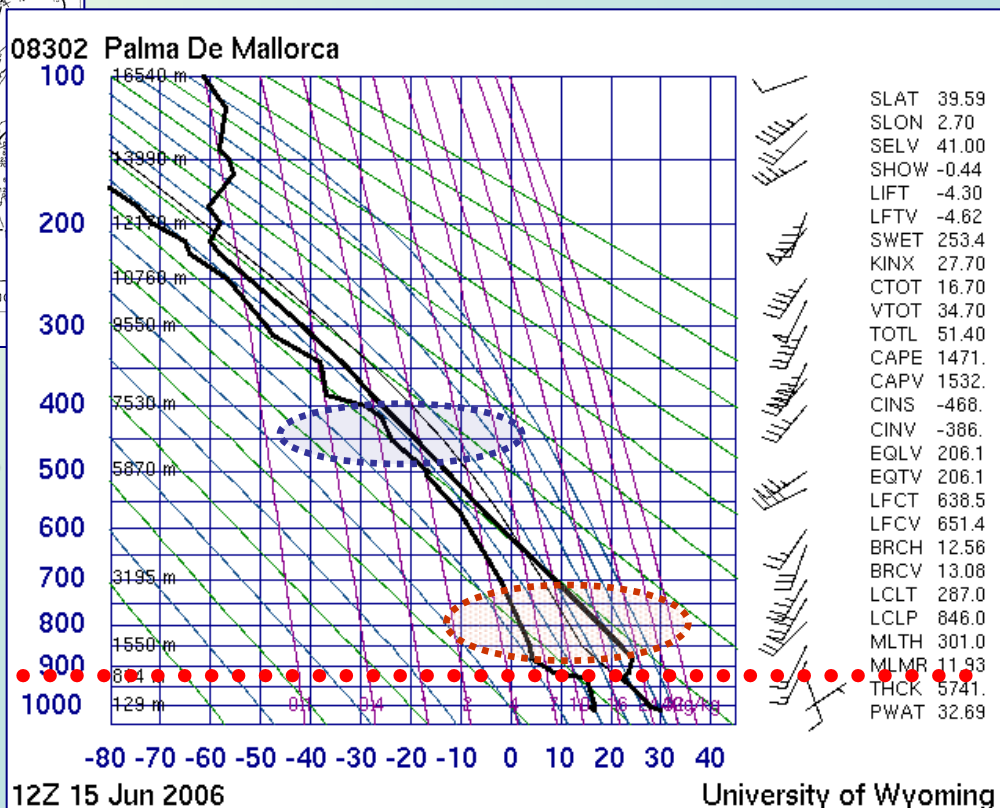


(1) low level Mediterranean air with weak surface depression

Synoptic characteristics



(2) warmer African air blowing around 850 hPa isobaric



(4) an instable layer between the African air and colder upper air

(3) an inversion layer separating (1) and (2)

rissaga warning system – Balearic Islands



Synoptic characteristics

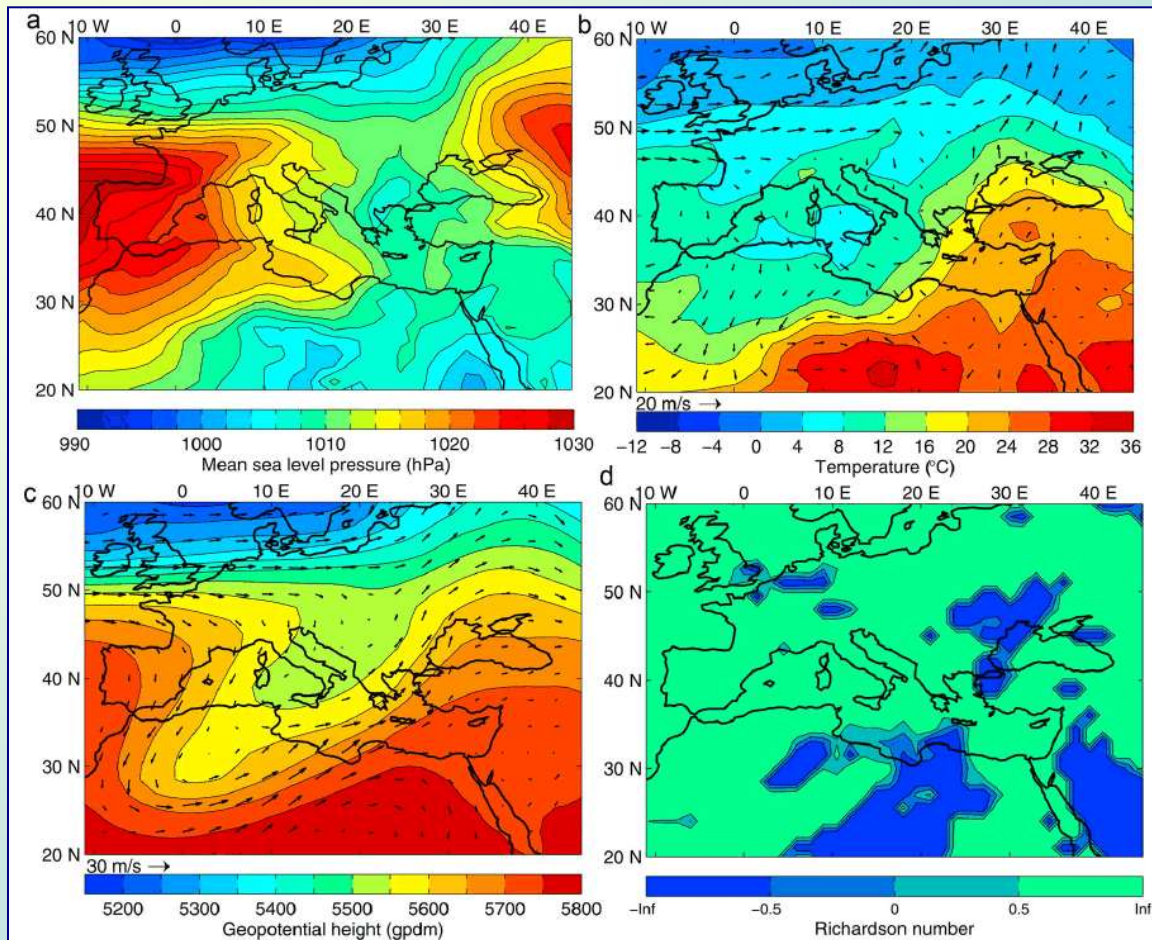


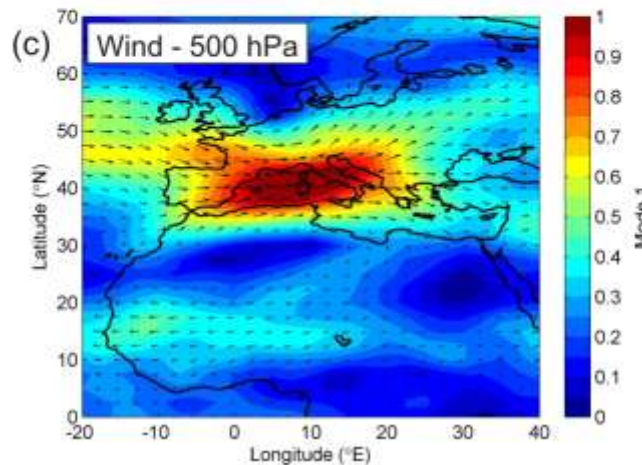
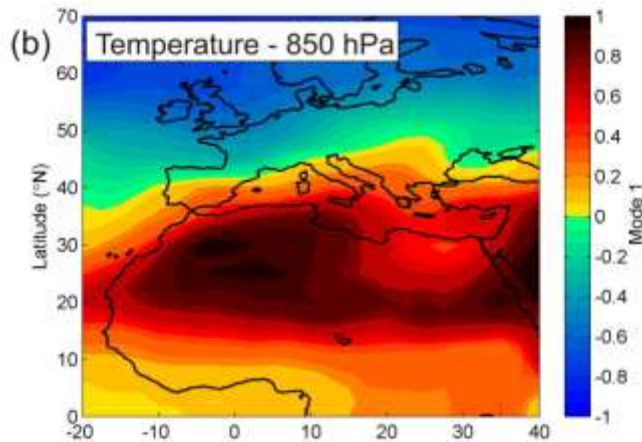
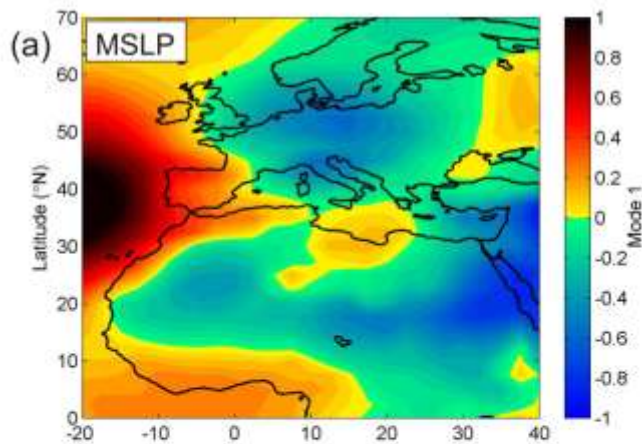
Figure 2. Synoptic charts as obtained from the ECMWF reanalysis fields of 7 May 2007 12:00 UTC: (a) surface air pressure, (b) temperature and winds at 850 hPa level, (c) geopotential height and winds at 500 hPa level, and (d) minimum Richardson number R_i at heights between 600 and 300 hPa.

Vilibić et al., JGR-O, 2010:

- synoptic setting for the Black Sea meteotsunami similar to the Adriatic and the Balearic meteotsunami settings



Synoptic characteristics

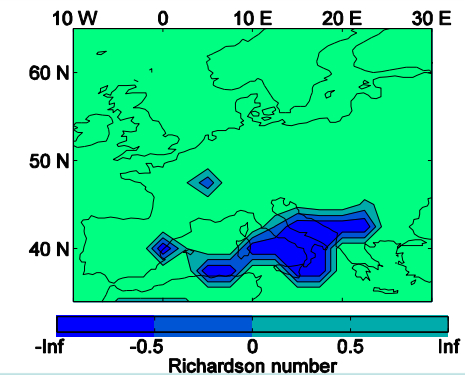
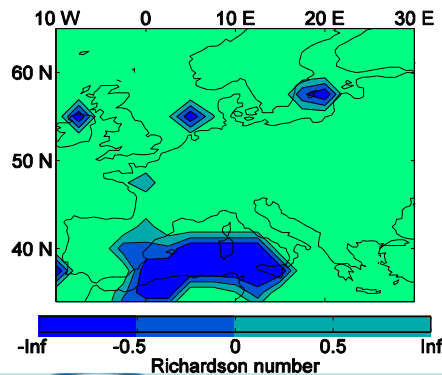
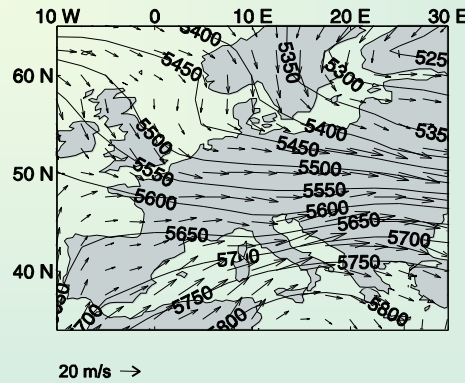
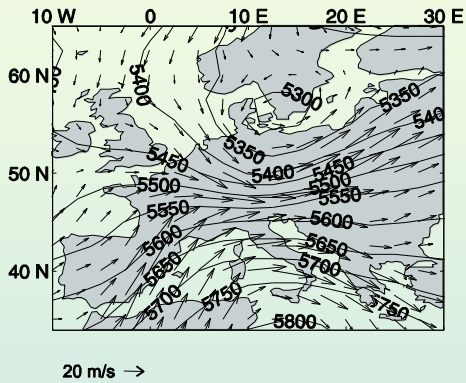
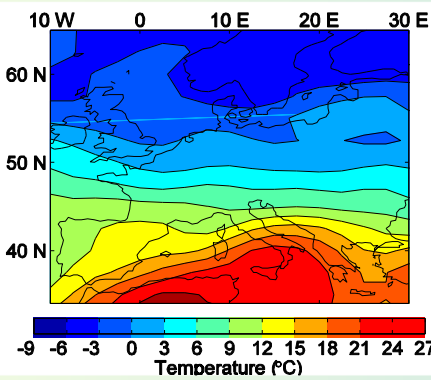
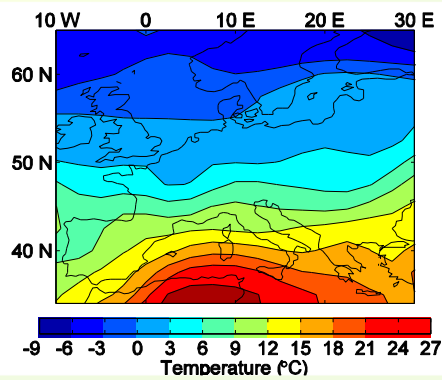


Šepić and Vilibić, JGR-O, submitted:

- first EOF of 16 strongest meteotsunamis measured at Rovinj (northern Adriatic)
- 99% SLP variability
- 89% 850-hPa T variability
- 45% 500-hPa wind variability



Synoptic characteristics



25 September 1990 (12 UTC)

26 September 1990 (12 UTC)

Šepić et al., PCE, 2009

- teleconnections between the Balearic and the Adriatic meteotsunamis due to synoptic settings
- out of the 32 “rissaga” events, 16 coincided with high-frequency sea level oscillations in Dubrovnik
- oscillations in Dubrovnik were more likely to happen if oscillations in Ciutadella were moderate or strong



- *Monitoring synoptic conditions*
- *Monitoring ground parameters (air pressure)*
- *Monitoring ocean parameters*
- *Real-time satellite observations*
- *Real-time meteo and ocean numerical models*



Monitoring ground parameters

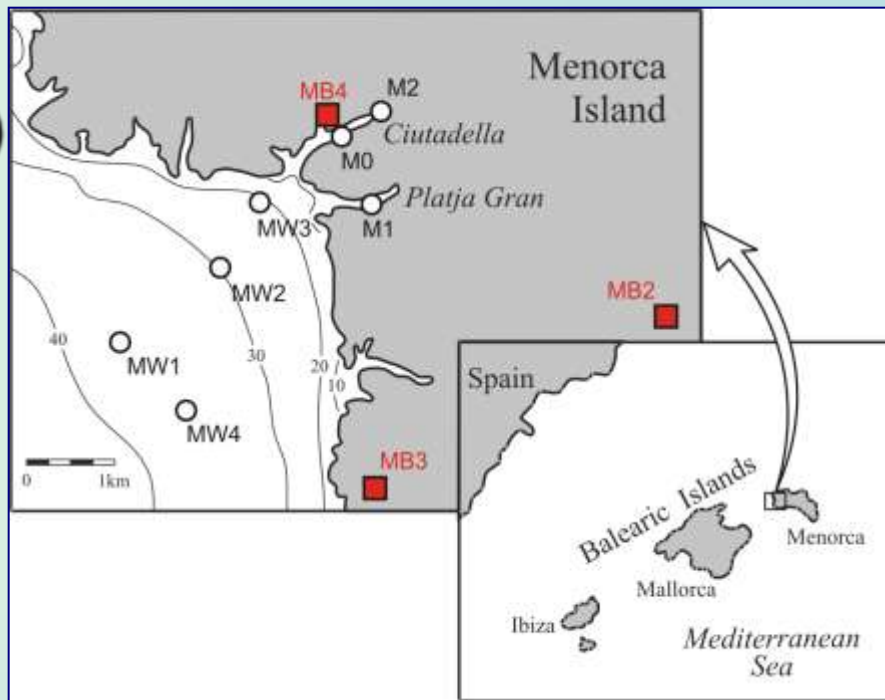
Operational possibilities

- Resolution max 1 min
- Real-time (latency max a few minutes)
- High-accuracy sensors (max 0.1 hPa or 1 mm)

Microbarographs



Coastal buoys



At least three triangles stations for determination of atmospheric disturbance speed, direction and dispersiveness (but located ahead of affected area)



Monitoring ground parameters

Operational possibilities



Pilot network of microbarographs (middle Adriatic Sea)

Such a network will allow for detection of potentially dangerous atmospheric disturbances 30-120 min prior to the occurrence of a meteotsunami

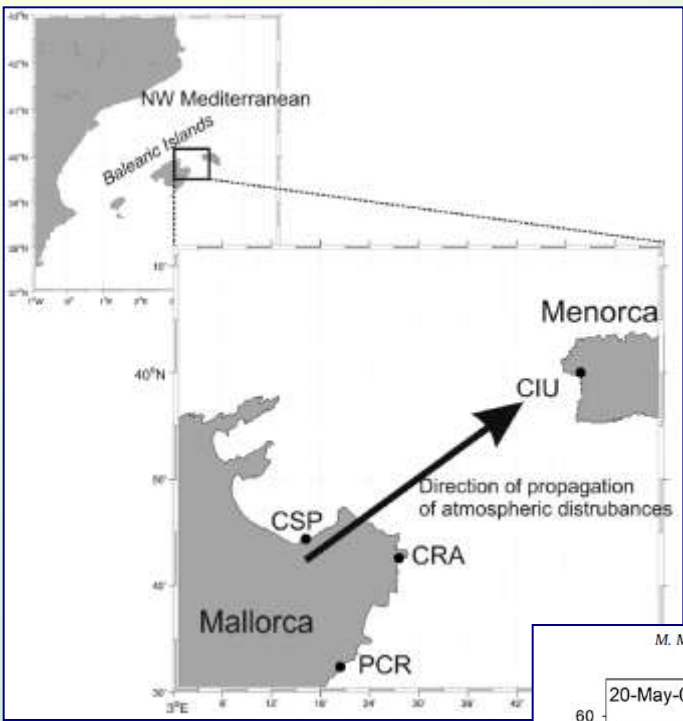


- hot spots
- microbarograph
- ★ open ocean buoy
- tide gauge
- existing tide gauge



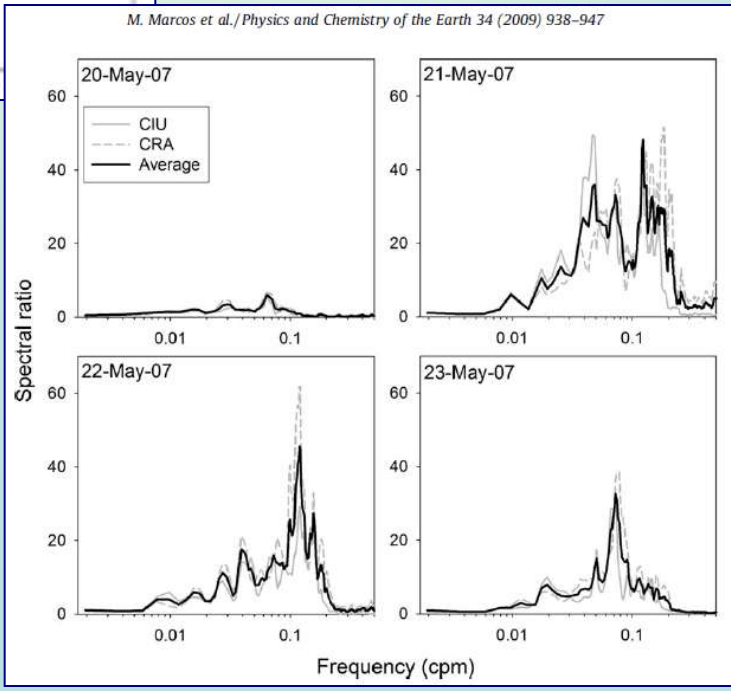
A proposal for early warning network (middle Adriatic Sea)





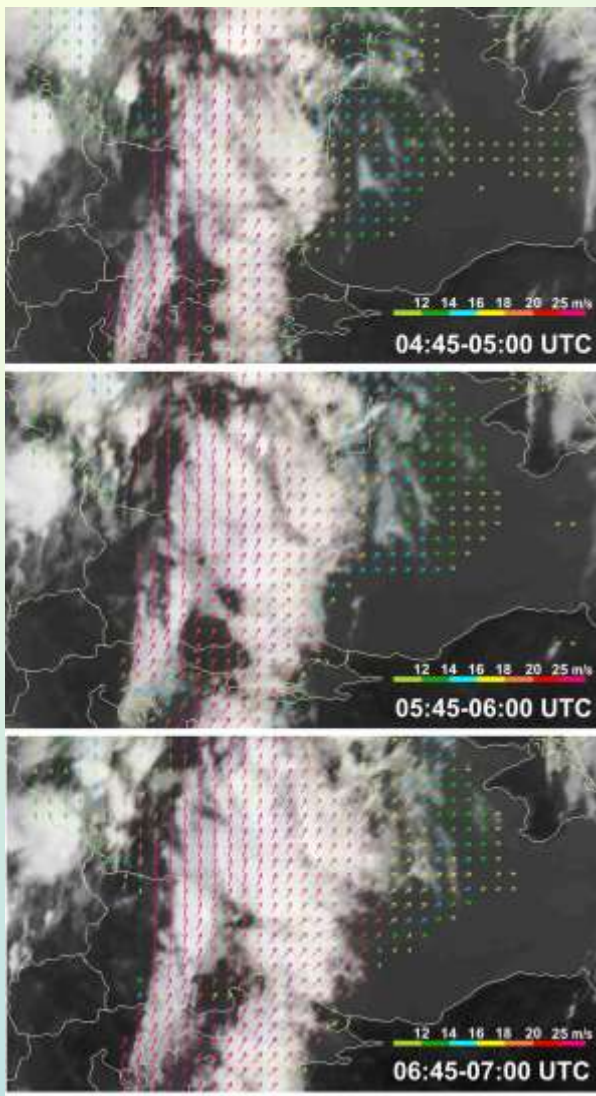
It could be possible to use *sea level oscillations observed at CRA to predict incoming rissaga events in Ciutadella Harbour*. These predictions could form part of a rissaga warning system designed to mitigate the damages inside the harbour. Similar approach probably could be used for some other regions of the Mediterranean where meteotsunamis occur, in particular for the eastern Adriatic: *certain 'beacon' sites can play the role of predictors for ports and harbours with probable destructive seiches.*

Marcos et al., PCE, 2009



Real-time remote observations

Operational possibilities



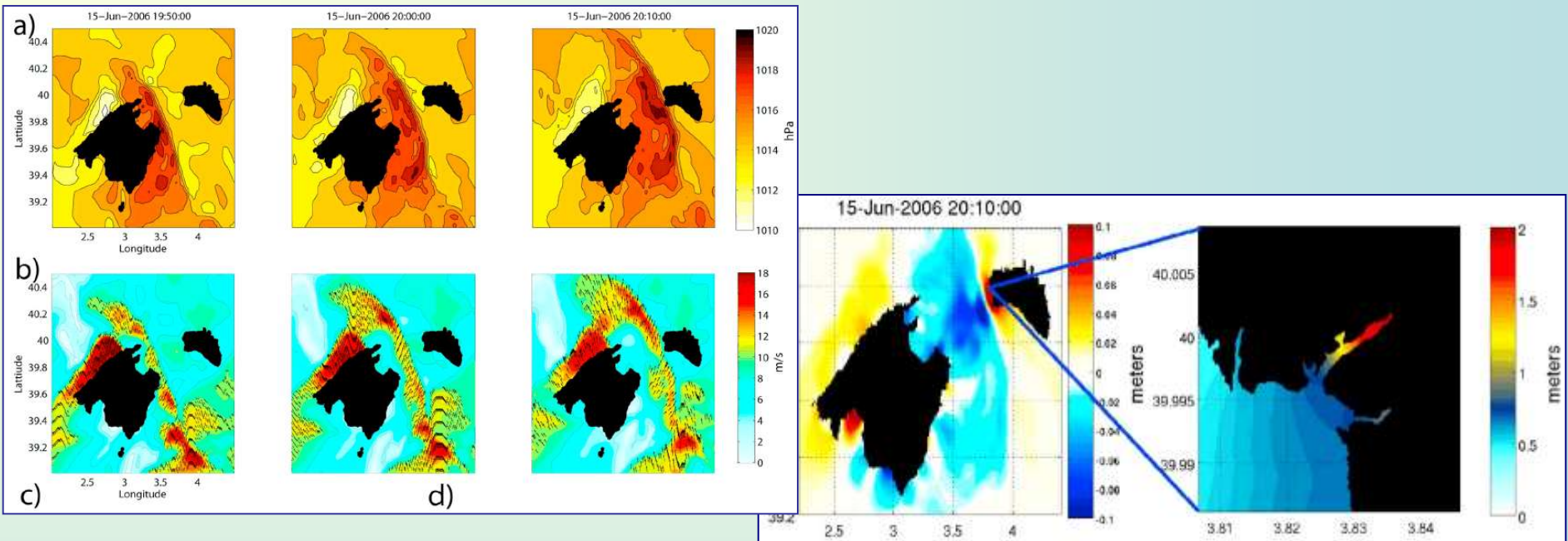
What about meteo radar network?



Estimation of cloud-top speed, and the connection to the base wind velocity



Real-time meteo and ocean numerical models **Operational possibilities**



Overall, the approach we considered here shows that **under certain atmospheric conditions meteo tsunamis** associated with travelling atmospheric waves and/or convective systems **can be forecasted**. It is important to note that the use of **standard meteorological forecast output at 3-hour intervals would not generate the meteo tsunami** in the ocean model. The pressure oscillation must be resolved to **about 2 minutes** in order to capture the sudden pressure change of order 0.3 hPa/min that appears to be required for sizeable Proudman resonance.



Renault et al., GRL, 2011



Instead of conclusions ...

- *There is a substantial progress on meteotsunami research*
- *Operational meteotsunami systems are at the beginning of its development*

